

MORICE BIOPHYSICAL STUDY

93 L/S.W.

**Wildlife Capability and Habitat
(Soils, Terrain, Climate and Vegetation)**



**Ministry of
Environment
and Parks**

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Victoria, B.C.

Wildlife Working Report No. WR-18
March 1986

Canadian Cataloguing in Publication Data

Main entry under title:

Morice biophysical study

(Wildlife working report, ISSN 0831-4330 ; WR-18)

Bibliography: p.

ISBN 0-7726-0595-5

1. Geology - British Columbia - Morice River Region.
2. Soils - British Columbia - Morice River Region.
3. Botany - British Columbia - Morice River Region.
4. Landforms - British Columbia - Morice River Region.
5. Habitat (Ecology) I. Fuhr, Brian, 1952- II. British Columbia. Wildlife Branch. III. Series.

HD319.B7M67 1986 333.7'097112 C87-092024-3

ACKNOWLEDGEMENTS

The author gratefully acknowledges the cooperation and assistance of Ministry of Forests Research Branch staff: Rick Trowbridge, Jim Pojar, Sybille Haeussler and Scott Lindeberg. Ministry of Environment staff Alan Edie, Rick Marshall, Gordon McKinnon and Ben van Drimmelen provided their assistance and information on the area. The Cartography section of the Surveys and Resource Mapping Branch drafted the maps, legends and report figures. Margaret Brighton and Darla Cooper did the word processing. To you, and those we may have unwillingly omitted, thanks!



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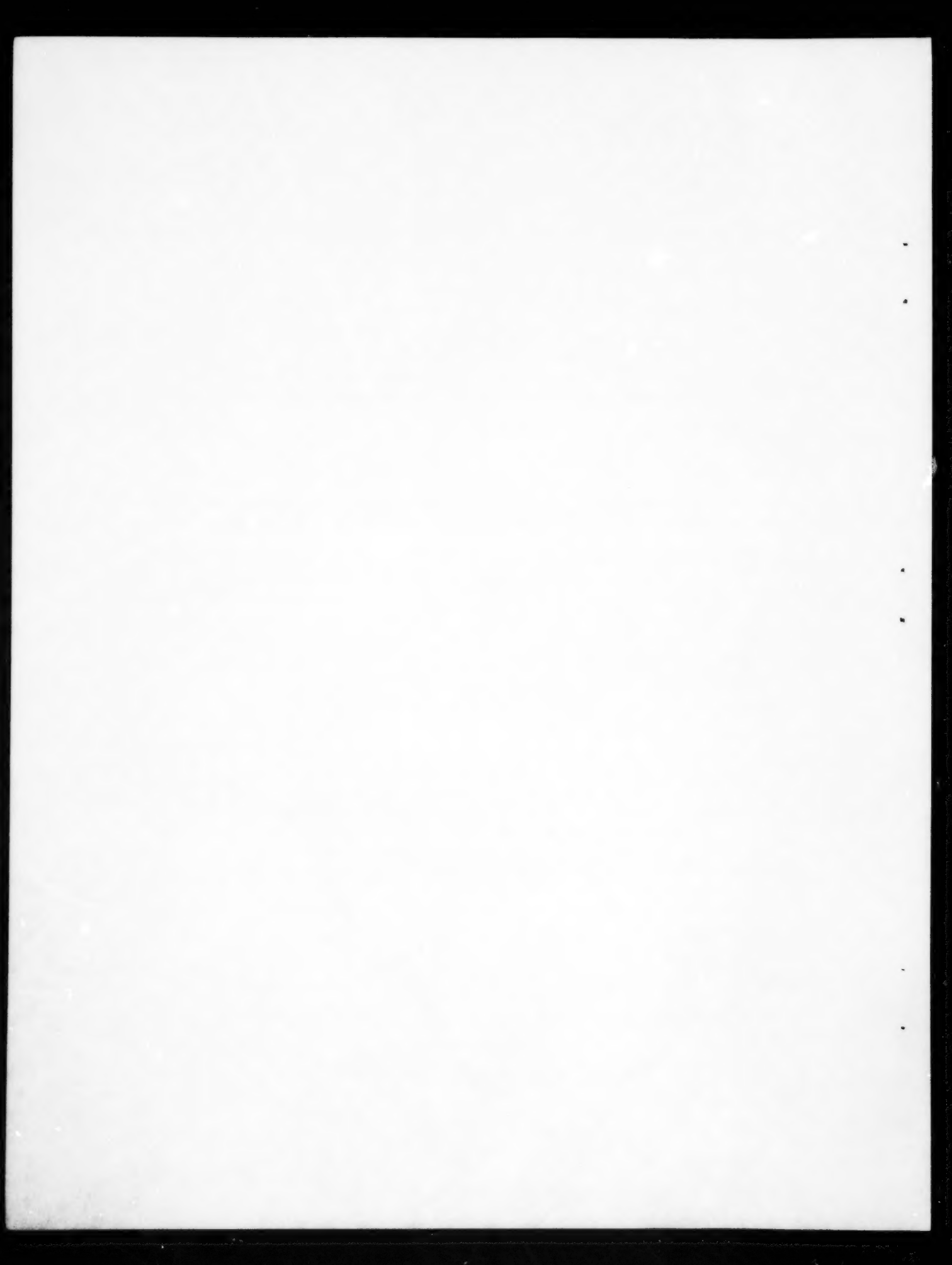
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CHAPTER 1. INTRODUCTION

The Morice Biophysical Study was initiated by the Planning and Assessment Branch of the B.C. Ministry of Environment. It is intended to provide reconnaissance biophysical information for this largely remote area. The area (see Figure 1) comprises four 1:50 000 scale NTS map areas: Lamprey Creek (93L/3), Corona Peak (93L/4), Burnie Lake (93L/5) and Thautl River (93L/6), covering an area of 3,690 square kilometers. In addition to this report there are several map products:

1. Terrain (4 maps)
2. Biophysical Soil Landscapes (93L/3 only)
3. Vegetation Landscapes of the Morice River Floodplain (portion of 93L/3 only)
4. Biophysical Ungulate Capability (4 maps)

The Lamprey Creek map area (93L/3) was studied in somewhat more detail as it contains areas of high capability moose winter habitat, it has considerable moose habitat/forestry interaction and has the greatest road access for sampling. Also, the proposed Kemano Completion hydroelectric development would potentially impact moose habitat along the Morice River floodplain, and was of particular concern while in the study area. The possible impacts of the Kemano Completion proposal are summarized by Ministry of Environment (1979), but do not include possible detrimental effects on moose habitat of the Morice River. Vegetation landscapes of the Morice River floodplain were done to help identify possible impacts within the study area. Areas downstream of this project were assessed for ungulate capability previously (Stewart, 1981).

The map products are of a reconnaissance nature, relying mainly on airphoto interpretation and a minimum of ground checking. They are applicable at the scale and level of survey at which they were produced. They provide a stratification of the materials, terrain hazards, soils and ungulate habitat capability at a scale of 1:50 000. Applications such as detailed road alignment or present habitat condition require further study.



Figure 1. Location of the study area.

The report follows the stepwise process of integrating the disciplines of climate, terrain, soils, vegetation and wildlife that is part of the biophysical method (Walmsley, et al., 1980). Each author provides details, background, methods and results.

CHAPTER 2. PHYSIOGRAPHY, BEDROCK GEOLOGY AND GEOLOGIC HISTORY

2.1 PHYSIOGRAPHY

The Lamprey Creek mapsheet (93L/3) and the southern portion of 93L/6 are described by Holland (1976) as being part of the Nechako Plateau physiographic region (Figure 2). This plateau area is characterized by relatively low relief and lack of dissection by major river systems. The majority of the study area is in the Bulkley Ranges of the Hazelton Mountains physiographic region. These mountains are characterized by ranges of relatively high relief, rugged, glacially sculptured peaks, separated by broad floored, U-shaped valleys. In the Hazelton Mountains, remnants of an older plateau surface are evident as relatively flat surfaced uplands surrounded by steep walled, glacially oversteepened U-shaped valleys. Plateau remnants east of the Burnie Lakes and Telkwa River provide rolling alpine terrain suitable for caribou range. The western margin of the study area is in the Coast Mountains physiographic region, an area of high relief, rugged, glacially sculptured peaks with large U-shaped, broad floored valleys between ranges.

2.2 BEDROCK GEOLOGY

Bedrock geology for this area was mapped by Carter and Kirkman (1969) and amalgamated into the geological atlas by Tipper et al., 1979.

The majority of the study area is underlain by Lower and possibly Middle Jurassic volcanic bedrocks including andesitic to rhyolitic tuffs, breccias and flows with lesser associated sedimentary rocks (Figure 3). These intermediate to acidic volcanic bedrocks tend to weather to form basic soils that are, in general, better soils for browse growth, than are acidic soils. Two relatively large belts of Upper Jurassic and lower Cretaceous sedimentary bedrocks (greywacke, siltstone, mudstone, conglomerate, with minor coal) occur in the McBride Lake and Holland Lake areas. These sedimentary bedrocks, being largely

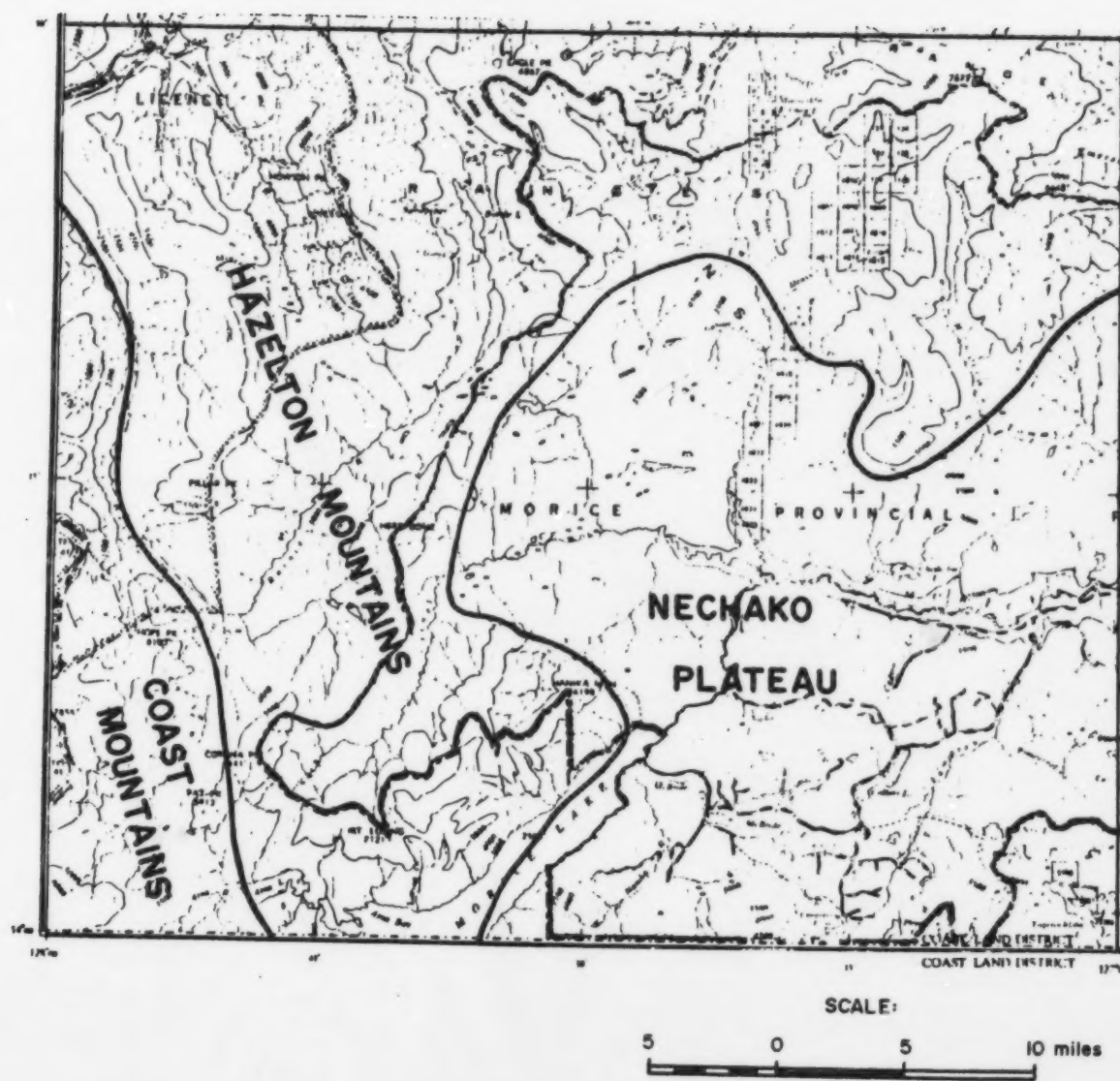


Figure 2. Physiographic regions.

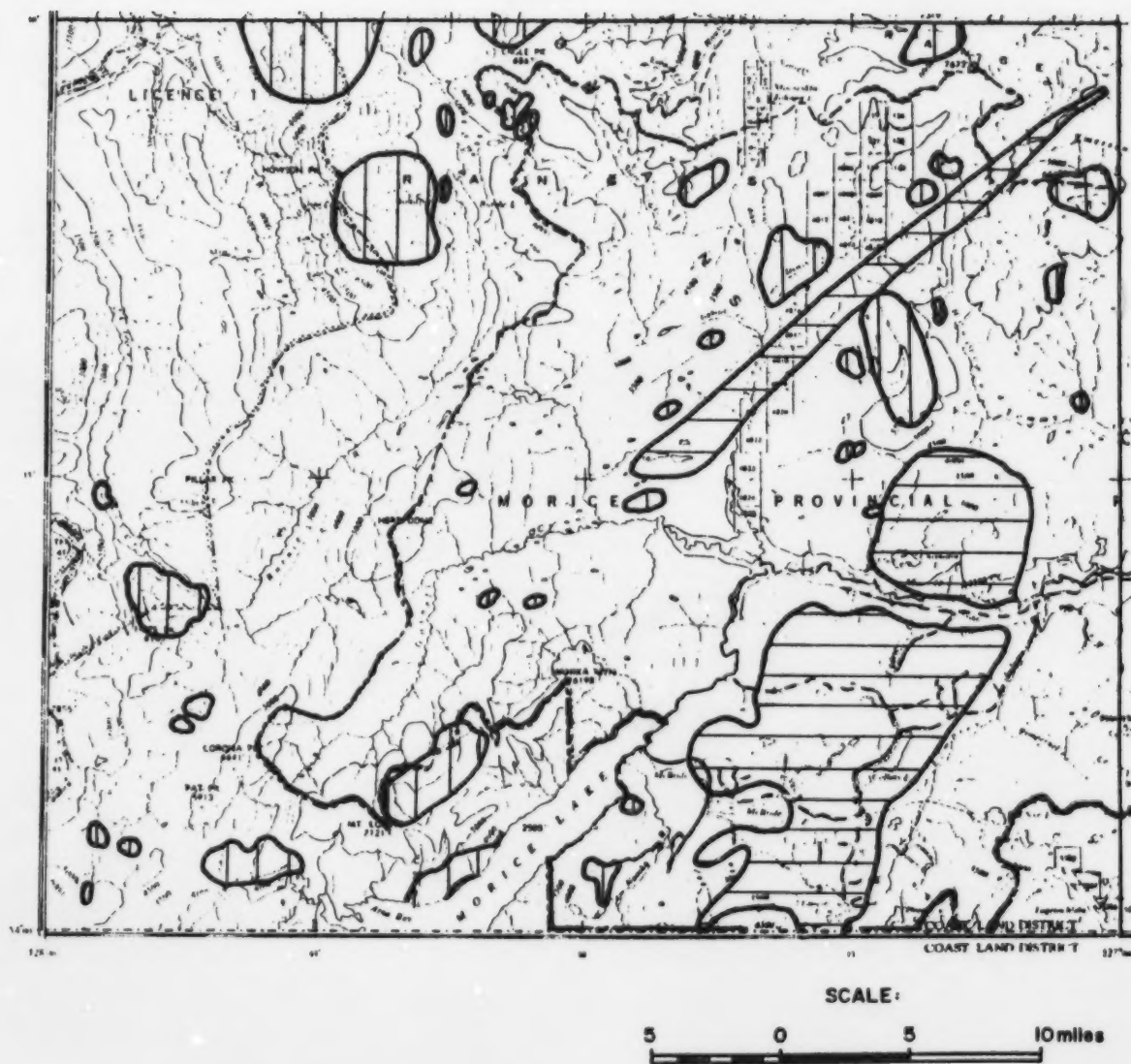


Figure 3. Bedrock geology (after Carter and Kirkham, 1969).

of volcanic origins, also tend to weather to produce basic soils. These older volcanic and sedimentary bedrocks and have been intruded by plutonic bedrocks (quartz monzonite, granodiorite, quartz diorite, with minor diorite and monzonite) in the Upper Cretaceous and early Tertiary. Subsequent erosion has exposed the plutons at a number of locations in the study area. These acidic to intermediate bedrocks tend to weather to produce acidic soils. At locations such as Howson Peak and Mount Loring, harder plutonic bedrocks often form steep cliff faces that provide good escape terrain for goats.

2.3 GEOLOGIC HISTORY

During the Tertiary geologic period a land surface of low relief was formed throughout the interior of the Province by a long period of erosion (Holland, 1976). A portion of this surface exists today as the portion of the Nechako Plateau in the study area (Figure 2). In the western and northern portions of the study area, the Tertiary land surface was re-elevated and deeply incised by stream and subsequent glacial erosion. Areas of high elevation, such as the Coast Mountains and the Howson Range have been sculptured by erosion to form a rugged mountainous topography lacking remnants of the ancient Tertiary plateau land surface. The Bulkley and Telkwa Ranges have been less severely eroded and consequently isolated segments of the ancient plateau surface occur at high elevations. The Nechako Plateau portion of the study area was not significantly uplifted, nor deeply dissected by stream nor glacial erosion, and has retained a land surface of comparatively low relief.

Quaternary deposits and landforms in the study area appear to have all been formed either during the most recent major glaciation, the Fraser Glaciation, or subsequent to it in the Recent Epoch. The initial ice accumulation of the Fraser Glaciation may have commenced in the Coastal Mountains and adjacent interior ranges as early as 29,000 years before present (Clague, 1981). However, for southern British Columbia at least, the interior plateau and plains appear to have remained largely ice-free until after 21,500 years B.P.

In the study area, the ice divide for east-west flow of Fraser Glaciation ice appears to have been along the ridge of the Howson Range and along a line from Pillar Mountain across the Clore Canyon to Dog's Ear Peaks (Figure 2). In the north-western portion of the study area, the orientation of major Coast Mountain and Bulkley Range valleys (Clore, Copper (Zymoetz) and Kitnayakwa), plus ample evidence of along valley glaciation (sculptured peaks and valley walls, as well as lineations) strongly suggest that as ice buildup increased in these ranges, flow was primarily westward out of the study area, and into the adjacent Kitimat-Terrace depression. Ice flowing south from source areas in the Howson Ranges eventually became sufficiently thick to override the adjacent highland plateau surfaces of the Nechako Plateau at Tom George Lakes and at Shea Lake and thus, to flow easterly. Ice moved out from the Coast Mountains on the western fringe of the study area and flowed in a north-easterly direction along the upper Clore and Atna valleys through the Bulkley Ranges and out into the lower elevation portions of the Nechako Plateau. Striae and lineations indicate that ice flow on the plateau was primarily to the east, as the Telkwa Ranges formed a barrier to the north. During this glacial period low relief plateau surfaces and portions of the floors and flanks of major mountain valleys were mantled by variable depths of basal till. Steep mountain slopes as well as hills and hummocks in major valley floors and on the plateau, were commonly glacially scoured and left either as barren rock, or as rock with interspersed pockets of till in depressions. The landform created by this glacial action is one of the factors contributing to the present distribution and quality of wildlife habitat in the study area.

During the late stages of the Fraser Glaciations, large quantities of fluvioglacial materials were deposited in front of and around stagnant ice, both on the plateau and in major mountain valleys. Ice downwasting appears to have been relatively rapid with no landforms suggestive of subsequent alpine/subalpine glacial activity being superimposed on Fraser Glaciation fluvioglacial landforms. However, alpine glaciation persists to the present in the highest ranges of the Howson and Coast Mountains.

In the Recent Epoch, rivers and stream in many valleys in the mountains have downcut into Fraser Glaciation drift, or bedrock, creating narrow inner valleys. However, on the plateau and in a number of major mountain valleys, stream gradients are low and extensive areas of fluvial sediments have accumulated. Recent Epoch mass wasting of glacially over steepened rock slopes, as well as till slopes, has left extensive lower slope areas buried under colluvium. The physical characteristics of glacial and post-glacial surficial materials, their landforms and present day geological processes are further discussed in Chapter 4.

CHAPTER 3. CLIMATE

3.1 CLIMATE

Physiographically the area is dominated by the southern Hazelton Mountains and parts of the Interior Plateau. Its location immediately east of the Coast Mountains has a great influence on the character of the climate of this area. Prevailing westerly winds flow over the region, most frequently bringing maritime Polar air over this part of the province. Much of the moisture is precipitated along the west facing slopes of the Coast Mountains as coastal air passes over the mountain barrier. Further drying occurs over the study area as the air subsides and warms while moving down the east facing slopes of the Coast Mountains. The result is a broad rainshadow over much of the Interior Plateau and a more discrete local rainshadow which exists as a 45 km wide band running parallel to the Coast Mountains and extending to a point 60 km east of the Coast Mountain Divide.

Precipitation is relatively light in this region. The rainshadow areas in the vicinity of Morice Lake (764 m) average 180 mm between May and September and 495 mm over the year (Air Studies Branch, 1980). The annual distribution of precipitation reflects the influence of coastal weather systems on the area since maximum monthly precipitation occurs in December and January. Despite this influence, the precipitation during these two months only accounts for 23% of the annual total. April is the driest month while precipitation is light and evenly distributed between June and November. About 45% of the annual precipitation occurs as snowfall.

Precipitation increases rapidly from east to west over the Hazelton Mountains, partially as a result of increasing elevation, but more the result of precipitation carry-over subsequent to the lifting of Pacific air masses over the Coast Mountains. The lower passes along the Coast Mountain Divide receive precipitation in excess of 1000 mm annually - more than 52% occurring as snowfall - and this total increases with increasing elevation.

Near to the western edge of the local rainshadow, snow depths reach an average maximum of 234 cm in April at an elevation of 1370 metres near Kidprice Lake. By June 1st, the snow depth averages 122 cm. Further east snow depths are considerably reduced and in the drier areas reach average maximum depths of an estimated 100 cm in April. Low elevation sites are snow free by mid-May.

The area is characterized by cool summers and moderately cold winters. July mean maximum temperatures average 20°C and mean minimums average 6°C. Temperatures and diurnal temperature range decrease with increasing elevation, although minimums are usually depressed in poorly drained sites. Winter temperatures are typified by January mean maximums of -12°C and mean minimums of -17°C. Throughout the winter and early spring, Arctic air frequently surges into the area from the east and stalls along the eastern slopes of the New Hazelton and Coast Mountains. This air mass is associated with clear skies and very low temperatures. Winter extremes in the area can exceed -40°C. The higher elevations of the New Hazelton mountains will be subjected to Arctic outbreaks less frequently. These data are from Meteorological Services, Waste Management Branch, Victoria, B.C.

CHAPTER 4. TERRAIN

4.1 INTRODUCTION

Of the four 1:50 000 scale mapsheets in the Morice study area, only the Lamprey Creek sheet (93L/3) has been surveyed for terrain at sampling intensity levels appropriate for conventional Terrestrial Studies Branch 1:50 000 scale reconnaissance mapping. The Corona Peak, Burnie Lake and Thautil River sheets (93L 4, 5, 6) which make up the more mountainous western and northern 3/4 of the study area have been surveyed at intensity level more appropriate for 1:100 000 or 1:250 000 scale terrain mapping, although the maps are presented at a 1:50 000 scale. These latter mapsheets have no road access and thus were surveyed less intensively.

4.2 PREVIOUS TERRAIN MAPPING

The surficial materials of the Morice area have been mapped by Gough (1974). However, for purposes of wildlife capability mapping, this earlier mapping was found to be deficient in characterization of geological processes (Section 4.5) and in the mapping of fluvial surficial materials. Thus, it was decided to update the 1974 mapping and to append soil drainage to the terrain map unit symbols (e.g. $\frac{Mvb}{(w)}$), to further aid wildlife capability mapping. Site and soil data from 17 Ministry of Forests research plots was also utilized in the re-mapping.

4.3 METHODS

The revised terrain mapping for this study area was carried out according to the 1983 revisions to the Terrain Classification System (unpublished, see Terrain Map Legend). This system identifies various types of surficial materials (unconsolidated deposits), provides data related to their texture and landforms, and describes geomorphic processes evident in the study area.

Initially the 1:63 360 scale black and white stereo air photographs used in the original terrain mapping were examined stereoscopically to

characterize surficial materials and geologic processes in terms of the 1983 Terrain Classification System. The map unit boundaries of the original mapping were retained, an exception being on floodplains where the original map units were often subdivided on the basis of fluvial process characterization.

The 1983 terrain and soil field survey program involved 5 1/2 days of helicopter traverses, 7 days of road traverses (93L/3), as well as 1 day of boat traverse along the Morice River. The objective of these traverses was to check designated terrain map units and symbols and to collect additional data on the physical characteristics of the surficial materials such as soil drainage and geological processes. Observations were recorded at 61 sites and 4 samples were collected for soil engineering analysis (grain size analysis, Atterberg Limits). Soils were also described at all sites and samples were collected for lab analysis at 48 sites in order to characterize major soil developments. Site, soil and lab analysis results are available from the British Columbia Soil Information System (Sondheim and Suttie, 1983).

Following the field survey and laboratory processing of samples, the terrain field and lab data, plus soil drainage data were compiled on 1:50 000 scale maps using the 1983 (unpublished) Terrain Classification System.

4.4 GENERAL CHARACTERISTICS OF SURFICIAL MATERIALS

This section describes the types of surficial materials found in the study area. It describes some of their significant physical properties, their local distribution and common landforms associated with each type of material.

4.4.1 Morainal Materials (Till)

Till is material transported beneath, beside, within and in front of a glacier. It is deposited directly from the glacier and not modified by any intermediate agent.

Till is the most common unconsolidated material on the Nechako Plateau portion of the Morice study area. Here it forms extensive mantles of varying thickness (e.g. Mvb, Mb, Mv), often over hummocky bedrock (e.g. $\frac{Mv}{Rh}$). In the major valleys of the Bulkley Mountains, such as the Kitnayakwa, relatively deep till deposits are also common on the valley floors with thinner tills on the slopes above. Some major mountain valleys such as the Burnie and upper Clore characteristically have extensive areas of blankets and veneers of till over hummocky to undulating bedrock surfaces (e.g. $\frac{Mvb}{Rh}$). Relatively high elevation plateau and upland surfaces such as those north of Morice Lake, north of Dockrill Creek and northeast of Burnie Lakes are commonly also characterized by till mantles, ridges and aprons of Neoglacial till (primarily terminal and lateral moraine) (e.g. $\frac{xMwra}{Rh}$).

Basal tills predominate in the study area and generally consist of moderately compact to compact, non-sorted, non-stratified materials with a wide range of particle sizes (pebbles, cobbles, boulders) in a matrix of sand, silt and clay. Till textures and clast lithology appear to be closely related to the predominant local volcanic and associated sedimentary bedrocks (see Section 2.2). These tills commonly have a silty sand to sandy silt matrix texture (e.g. $\$sM$) (Table 1), with very variable percentages of coarse fragments (primarily pebble and cobble sized fragments of volcanic rock).

The basal tills generally consist of a weathered upper stratum that is relatively loose, non-compact and relatively permeable (rapid to moderately well soil drainage depending upon slope position). The unweathered (basal) portion of the till is compact, relatively impermeable and restricts drainage. On toe-slopes and in depressions, water accumulation above the basal till commonly results in localized wetter habitats that usually have an increased diversity and abundance of browse species. In comparison to coarse textured colluvial and fluvioglacial deposits, all study area tills examined have more favourable soil moisture retention properties and consequently are usually capable of supporting a greater diversity of browse species.

Unweathered basal till is relatively resistant to shear and deformation due to preconsolidation by glacial ice. However, the few samples of basal

Table 1. Textural Analysis and Atterberg Limits for Morice Study Area T111 Samples

SAMPLE NUMBER*	PERCENTAGE BY WEIGHT				ATTERBERG LIMITS	
	Coarse Fragments (>2 mm)**	Sand (2- .0625 mm)	Silt (.06254- .002 mm)	Clay (<.002 mm)	Liquid Limit	Plastic Limit
82-17752	71	13	10	6	20.36	15.27
82-17757	43	25	24	8	22.83	19.53
82-17782	72	15	10	3	15.13	11.56
82-17783	14	34	35	17	22.03	15.97

* For locations see terrain map.

** % by wt. coarse fragments - visual estimate in field, % sand, by weight of silt and clay - laboratory analysis.

tills suggest that they are generally weakly plastic (liquid limits of 15 to 23, Table 1). Thus when wetted they may be subject to deformation and shear failure, especially on large cut faces. They are not as suitable as fluvioglacial materials for road beds, nor for operation of heavy machinery during wet weather.

Tills with coarser textures may occur in areas dominated by coarse drained plutonic bedrock outcroppings, but were not observed during the very limited number of stops in such areas during the field survey. Such tills would have more favourable roadbuilding and machinery operation characteristics, but less favourable with regards to the maintenance of browse species, being more droughty and acidic.

Ablation till (till deposited as a result of downwasting of glaciers) was not observed in the mountainous portions of the study area. However, shallow mantles of relatively coarse textured ablation till commonly veneer finer textured basal tills on the Nechako Plateau. By hand texture and visual estimates ablation till has a sandy gravel texture and is loose and permeable (rapid to well drained). It resembles fluvioglacial gravels in physical characteristics and in ability to support browse species.

Neoglacial tills commonly consist primarily of angular coarse fragments (e.g. xM) (based on visual and hand texture estimates) with lesser and variable amounts of interstitial sand, silt and clay. These relatively young, coarser textured tills are commonly relatively loose, non-compact and permeable (rapid to well drained). They are commonly sparsely vegetated due to their location in and near high elevation cirques and their coarse textures.

4.4.2 Colluvial Materials

Colluvial materials are products of mass wastage, chiefly consisting of rubbly and blocky bedrock material. They are perhaps the second most common surficial material in the study area, especially in mountainous topography. Here, they are common as aprons and fans (e.g. rCa, rCf) at the bases of steep mountain slopes, as well as blanketing and veneering portions of valley walls (e.g. rCvb). On the hummocky or undulating bedrock controlled

topography commonly seen on the Nechako Plateau and on the floors of major valleys in the mountains, colluvium forms extensive mantles of rubbly materials (e.g. $\frac{rCvb}{Rhr_u}$), weathered largely in situ. Here it is commonly mapped in association with veneers of weathered till (which it often closely resembles) and hummocks of rock. Higher elevation plateau/upland surfaces are also partially mantled by very rubbly colluvium weathered in situ, largely by frost shattering.

The colluvial materials of the study area were not analyzed for textural analysis nor for Atterberg Limits. However, hand texture and visual estimates suggest that a rubbly texture (dominantly <256 mm rubble, with lesser blocks as well as lesser interstitial sand, silt and clay) is appropriate, particularly in the mountains. The rubbly colluvial materials are relatively loose and permeable (generally rapid to well drained), creating droughty sites with a relatively sparse vegetative cover. However, the toes of colluvial fans and aprons are commonly moister and are more lushly vegetated. Also, morainally derived colluvium on the Nechako Plateau may have a finer texture and support better plant growth. Rubbly colluvium is not usually subject to plastic behavior and except for slope limitations, is suitable as road subgrade and as a bearing surface for heavy machinery operation.

Colluvium derived from failed tills has physical properties intermediate between colluvium and till and would generally be subject to plastic behaviour.

4.4.3 Fluvio-glacial Materials

Fluvio-glacial materials are fluvial materials deposited in contact with glacier ice. They commonly show evidence of ice melting such as irregular topography, irregular bedding or slump structures (e.g. F_{hd}^G). In the study area, fluvial materials deposited in front of glaciers (outwash plains) have also been mapped as fluvio-glacial deposits (e.g. F_t^G).

Fluvio-glacial materials are common in major valleys in the mountains (e.g. the Howson - Thautil Creek Valley), due to large volumes of water from

active/stagnant mountain valley glaciers. These mountain valley deposits include kame terraces and fluvioglacial fans both deposited adjacent to, or partially on ice, as well as outwash plains on broad valley floors. These deposits have often been dissected by post glacial stream downcutting and today are evident as terraces above incised inner valleys, as in the Clore River Valley.

On the Nechako Plateau, downwasting of stagnant/dead ice masses resulted in development of extensive fluvioglacial outwash plains generally with associated kame terraces with kettles, eskers and hummocky ablation till. Plateau fluvioglacial deposits have not generally been deeply dissected by post glacial stream downcutting. Most glacial meltwater channels on the plateau are not deeply entrenched and have come to be occupied by small, irregular meandering misfit streams.

Fluvioglacial materials in the study area were not sampled for mechanical analysis. Based on visual estimates and hand texturing, they generally have a sandy gravel texture (usually >50% pebbles, cobbles and lesser boulders with <50% sand). In valleys such as Houston Tommy that are characterized by extensive outwash plains, sandy textured fluvioglacial deposits occur on the floors of some meltwater channels.

The fluvioglacial materials are non-plastic and non-consolidated and due to their coarse-textures have good bearing capacities. They are highly permeable (rapid to well drained) and consequently are dry sites with a relatively sparse browse cover. The terrestrial lichens commonly used by caribou may be abundant on these sites.

4.4.4 Fluvial Materials

Fluvial materials are those transported and deposited by streams and rivers (alluvial materials). In the Morice study area the most extensive deposits of fluvial materials are low gradient floodplains associated with irregular meandering creeks and rivers on the plateau and in major mountain valleys, such as the Atna and Clore - Burnie valleys (e.g. sgF_p^A-I). As previously mentioned, these water courses are usually misfit streams in larger meltwater channels. They are flanked by the escarpments of higher

fluvioglacial terraces and fans, and occasionally by the lower escarpments of younger fluvial terraces. Portions of these floodplains are often occupied by active fluvial fans and organic deposits. Smaller mountain valleys with steeper gradients generally have incised streams with very small floodplains.

Visual and hand texture estimates of fluvial materials indicate that they have a sandy gravel texture (>50% pebbles, cobbles and boulders, <50% sand and a few percent silt and clay) (e.g. sgF^A_P , sgF^A_f , sgF_t). On the plateau, floodplains tend to have an overlying silty or sandy surface strata that in turn may be partially capped by organic deposits (e.g. $Op = F^A_P - I$).

The coarser textured fluvial materials are non-consolidated, loose, non-plastic and permeable. They have favourable characteristics for road building and heavy machinery operation except for seasonal limitations due to high water tables and flooding. Silty textured floodplains are somewhat plastic, much less permeable, and much less suited for road building materials and heavy equipment operation. Floodplains are commonly mapped as having imperfect soil drainage, but often have areas of poor to very poor soil drainage with organic cappings. Floodplain sites are highly productive of browse species. Associated inactive fluvial terraces and fans may, however, be well to rapidly drained and have engineering properties and vegetation characteristics similar to those for droughty fluvioglacial materials.

4.4.5 Organic Materials

Organic materials result from the accumulation and partial decay of vegetative matter. The organic materials in the study area most commonly occur as flat surfaced deposits over 1 m in depth, along low gradient floodplains (e.g. $Op = F^A_P$).

The organic materials most commonly consist of mesic textured (partially decomposed) sedges or mosses. They are very compressible and have very low bearing strengths and consequently are not suitable for road building nor

heavy equipment operations. The water table is generally near, at, or above the surface (very poor soil drainage).

4.4.6 Glaciolacustrine Materials

Glaciolacustrine materials are those deposited in direct contact with glacier ice, or those containing evidence of deposition in close proximity to glacier ice. Glaciolacustrine materials were seen only in sections on the Clore River where they are overlain by till (e.g. $\frac{Mb}{CSL}$ G) and in one small area downstream of the Lamprey Creek - Morice River confluence. Hand texture estimates indicate that these latter materials have a sandy silt texture (>50% silt, <50% sand with lesser clay). The glaciolacustrine materials were not sampled for mechanical analysis but appear to be non-consolidated, moderately compact, weakly plastic and relatively impermeable. They have significant limitations for road building or heavy equipment operation, especially when wet. They also have favourable moisture retention characteristics and depending upon topographic location and soil drainage, are similar to tills in ability to support browse species.

4.4.7 Ice

Ice includes areas of snow and ice where evidence of active glacier movement is present. The only extensive areas of ice in the study area are the glacier fields on the Howson Range and in the Coast Mountains on the western edge of the study area.

4.5 GEOLOGICAL PROCESSES

Geological processes are processes that are currently modifying, or have modified surficial materials and surface expressions. A number of hazardous geological processes are evident in the mountainous portions of the study area and include failures in: bedrock (rockfall, rockslide, bedrock slump), surficial materials (debris torrent), and in ice and snow (avalanches). High elevation peaks, uplands and plateau surfaces are also subject to a number of much less hazardous, but significant periglacial processes

including cryoturbation, nivation and solifluction. Fluvial processes evident in the study area include development of irregular sinuous, meandering and braided watercourses. Gully erosion is a significant process on steep mountain slopes, on till slopes, on fluvial fans, and in canyons. These processes influence the type and quality of wildlife habitat.

4.5.1 Rockfall

Rockfall is the detachment of masses disintegrating bedrock from steep slopes and their descent mostly through the air by leaping, bouncing and rolling. In the Morice study area, it is also used to include small rock-slides and rock avalanches that contribute to talus accumulation. It generally involves repeated falls of small amounts of debris resulting in fresh scars on rock faces (e.g. Rs-Rb) and fresh talus debris (e.g. rCaf-Rb). It is most common in steep, rocky mountain summit areas, but also occurs to a limited extent on steep rock faces at lower elevations. Very active areas are poorly vegetated and present a hazard to mountain goats traversing these areas.

4.5.2 Rockslide

Rockslides are the rapid downslope movement of masses of bedrock along well defined slip planes (e.g. Rs-Rr). The moving mass disintegrates or is severely deformed. This process was observed only in the Kitnayakwa River-Steward Creek valleys where numerous small, and one extensive, deep seated rockslide failures have occurred in volcanic bedrock.

4.5.3 Bedrock Slumps

Bedrock slumps are the sliding of internally cohesive masses of bedrock along a surface of rupture that is either concave upwards, planar or irregular. This process was observed only in the Kitnayakwa River and Steward Creek valleys where it is characterized by scarp slopes (slide scars) and tension cracks associated with displaced slump blocks in volcanic rock (e.g. Rs-Fm).

4.5.4 Debris Torrents

Debris torrents are the rapid flow of slurries containing rock fragments and vegetative debris down well defined, steep mountain side stream channels. In the Morice study area this process is identified by debris accumulation, sediment plugs and/or scattered large blocks in the channels carved in the mountain slopes (e.g. $rCv/Rs-Rt$), and by coarse, non-sorted debris, large blocks, levees and wood debris on fans affected by debris torrents (e.g. $rCf-Rt$). This process is common on steep, gullied slopes in the Coast Mountains, but also occurs less frequently on similar topography in the Bulkley Ranges.

4.5.5 Avalanching

Avalanching is the rapid downslope movements of snow, ice and other incorporated debris on steep mountain slopes and in gullies (e.g. $rCv/Mv-A$), often extending out onto valley floors (e.g. $rCa-A$). It is prevalent throughout the mountainous portions of the study area. Frequent avalanches may preclude winter range use by mountain goat on otherwise suitable range and may be an important mortality factor.

4.5.6 Cryoturbation

Cryoturbation consists of heaving, churning and related movements that result from repeated freezing and thawing of moist unconsolidated sediments. In the Morice study area it is characterized by development of sorted circles, stone stripes, and earth hummocks on alpine/subalpine slopes above 1500 m and on all aspects (e.g. $rCv-C$). It also commonly occurs on plateau surfaces above 1500 m (e.g. $\frac{Mv}{Rhu}-C$).

4.5.7 Nivation

Nivation is the erosion and enlargement of hollows containing snow patches due to a combination of freeze-thaw processes, physical and chemical effects of meltwater, solifluction and soil creep. In the Morice study area it is characterized by circular to elongate "snow patch hollows" on alpine/

subalpine slopes (Rhs-N) and on plateau surfaces (e.g. Mv/Rhu-N). It is most common on "northerly" aspect slopes at elevations greater than 1500 m. Such areas are covered with deep snow until summer.

4.5.8 Solifluction

Solifluction is the slow downslope movement of moist or saturated, seasonally frozen unconsolidated material over a relatively impermeable substrate that may be bedrock or frozen ground. In the Morice study area it is characterized by lobes, sheets or terraces of soliflucted materials on alpine/subalpine slopes (e.g. rCvb-S). It generally occurs above 1500 m on all aspects.

4.5.9 Irregularly Sinuous Channel

The irregularly sinuous channel fluvial process is evident where the channels display irregular turns and bends without repetition of similar features (e.g. sgF^Apt-1). It is the typical fluvial process on relatively low gradient, laterally confined floodplains throughout the Morice study area. Floodplains subject to this process may be seasonally inundated by slow to rapidly flowing overbank flooding. Channel bank erosion and channel shifting may occur.

4.5.10 Meandering Channel

The meandering channel fluvial process occurs where a channel is characterized by a repeated pattern of bends with uniform amplitude and wave length (e.g. sF^Ap-M). This process was mapped on low gradient reaches of the upper Clore River, Gosnell Creek and the Morice River. Areas subject to this process are generally seasonally inundated by slow moving overbank flooding. Gradual channel bank shifting and channel bank erosion occur. This is one of the most important processes in renewing seral progression on moose range in the area.

4.5.11 Braiding Channel

The braiding channel fluvial process occurs where the active channel zone is occupied by many diverging and converging channels separated by bars. Braiding channels are subject to seasonal flooding by rapidly moving water, channel and bar shifting, lateral instability and abrupt channel shifts (e.g. sgF^Ap-B). It is a typical process on fans and on the floodplains in large, relatively low gradient mountain valleys that are receiving excess sediment from steep tributary valleys and/or glaciers. Such areas provide some good summer moose range, but are often poorly vegetated.

4.5.12 Gullying

The gullying process is evident where surfaces have been eroded into long, narrow, steep sided depressions by running water, mass movement processes, or snow avalanches. In the Morice study area this process is commonly observed on steep rocky mountain slopes (e.g. Rs/rCv-V), on till slopes (e.g. Mvb-V) and in canyons (e.g. Us-V). Flooding, bank erosion, debris torrents and snow avalanching are typical hazards encountered in gullies. The complex topography created by this process often results in mosaic of feeding and secure resting sites for ungulates.

CHAPTER 5. BIOPHYSICAL SOIL LANDSCAPES OF THE LAMPREY CREEK AREA, 93L/3

5.1 INTRODUCTION

The Lamprey Creek 1:50 000 mapsheet 93L/3 is located between 54°00' and 54°15' north latitude and 127°00' and 127°30' west longitude. This map area covers 91,500 ha in the southeast quarter of the study area.

There are no settlements within the Lamprey Creek map area, however, considerable road access has been developed to access forest resources and one high elevation mining exploration area in the southwest corner of the study area. There are two B.C. Ministry of Forests low maintenance picnic sites, one on Morice Lake and another at the mouth of Lamprey Creek, as well as a cabin on McBride Lake. Major recreational opportunities for boating occur on McBride Lake, Lamprey Lake, and Tagetochlain Lakes as well as on the Morice River. Sport fishing has long been of major importance and the value of the Morice River, Nanika River and Morice Lake waters for salmon are detailed by Shepherd, 1979.

Most of the bedrock outcrops within the Lamprey Creek map area are fine grained and of volcanic and sedimentary origin. Coarse grained acid intrusive rock outcrops are limited to an area west of the Nanika River and south of McBride Lake. These two bedrock groups were recognized within the biophysical soil landscape legend. Table 2 shows how the bedrock geology maps have been generalized.

There are 3 biogeoclimatic zones present within the Lamprey Creek area; the Alpine-Tundra, Engelmann Spruce Subalpine Fir and Sub-boreal Spruce Zones (Ministry of Forests, 1984). Most of the area is within the Subalpine Fir Subzone of the Subboreal Spruce Zone. The Engelmann Spruce Subalpine Fir Zone is restricted to areas above 1070 m (3500') such as Pimpernel Mountain and the southwestern edges of the map area. The Alpine-Tundra zone only occurs northwest of Morice Lake. A very small area, the lowest elevation within the study area, 670 m to 730 m (2200' to 2400') along the Morice River is within the Spruce Subzone of the Subboreal Spruce Biogeoclimatic Zone. This area has the lowest snowfall and longest

Table 2. Generalized Bedrock Groupings for the Lamprey Creek Map Area* 93/L-3

1. Fine grained sedimentary and volcanic

O	Endako Group	andesite, basalt, dacite
KRR	Red Rose	shale, greywackie, conglomerate, coal gabbro
ETb		rhyolite, dacite, trachyte,
KTOL	Oosta Lake Group	sandstone, shale conglomerate
IJTN	Telkwa, Nilkitkwa Group	basalt, andensite, breccia, tuff, shale, siltstone
KBB	Brian Boru	andesite to rhyolitic tuff, breccia, flows

2. Coarse grained acidic intrusive

ETg	Early Tertiary	quartz monzonite, granodiorite, quartz diorite
EJg	Topley Intrusions	quartz mozonite, granodiorite

* after Tipper H.W., R.B. Campbell, G.C. Taylor and D.F. Scott, 1979, Parsnip River British Columbia. Sheet 93. 1:1 000 000 Geological Atlas. Geological Survey of Canada. Energy, Mines and Resources.

additional information on bedrock geology from:

Carter, N.C. and R.V. Kirkham, 1969. Geological Compilation Map of the Smithers, Hazelton and Terrace Areas. British Columbia Department of Mines and Petroleum Resources. Map 69-1.

growing season. These zones and subzones provide a level of stratification of soils below physiographic region and bedrock type.

The summary of the ten short term climate stations within the Lamprey Creek map area appears in Chapter 3.

5.2 PREVIOUS SOIL AND TERRAIN WORK

Five site, soil and vegetation plots were collected by the B.C. Ministry of Forests in support of biogeoclimatic classification. Their locations are marked on the map.

The 1:50 000 scale landform map (93L/3) produced by N. Gough in 1974 was used as a base and revised to the unpublished Ministry of Environment Terrain Classification System. Process modifiers such as flooding, and surface expressions such as depth were the major additions to Gough's maps. The eight detailed soil profiles sampled during this earlier project supplemented the soil data collected during this study.

B.C. Ministry of Forests carried out detailed (1:20 000 scale) soils mapping in an area south of Pimpernel Creek and north of Bill Nye Lake (Lindebergh and Trowbridge, 1984). The general descriptions of their soil units were used to help extrapolate to the surrounding areas with similar materials and drainage conditions. No soil samples were analysed from this more detailed study area.

Soil samples collected by L. Lacelle in 1983 for this project were also used if they were collected within a biogeoclimatic subzone common to the Lamprey Creek map area.

There are published soil maps to the east of the study area (Runka, 1972). Although many of the soils recognized in the Lamprey Creek area have soil association names in the adjacent map area 93/L 2, no attempt was made to correlate with these association names.

5.3 METHODS

Prior to field work, aerial photographs were studied to gain familiarity with general tone, pattern and texture of the photographs so as to relate these to identification of drainage conditions and material types when in the area. The aerial photographs were taken at a 1:50 000 scale in 1979; flight line BC 79076 photo numbers; 154-170, and 242-249.

A total of eight days were spent in the Lamprey Creek map area, 93L/3, in July, 1983. Most roads were travelled by vehicle, supplemented with some short hikes to check map units. Boat access was used along the Morice River and 1 1/2 hours of helicopter time was used to access more remote areas.

Thirty-three plots including site, soil, vegetation and wildlife descriptions were collected in accordance with the standards set out in Walmsley et al., 1980. The site, soil and laboratory analysis is stored in the B.C. Soil Information System as described by Sondheim and Suttie, 1983. Additional information collected in the field was recorded on aerial photographs and incorporated into the terrain and biophysical soil landscape maps.

Appendix A shows how the field samples are located within the biophysical soil landscape legend. This is intended as a rough guide to the reliability of physical and chemical descriptions. Usually, landscape types with more described sample sites are better understood.

Laboratory analyses include: soil reaction (pH in calcium chloride), organic matter content (organic carbon), total nitrogen, cation exchange capacity, exchangeable cations, extractable iron and aluminum percentages. Some physical analyses were carried out to determine sand, silt and clay percentages for selected mineral horizons. Organic horizons were analysed for reaction, rubbed fiber content and pyrophosphate index.

Since the sampling was selective, the summary of the data relies on the experience of the surveyor. Some plots are considered to better

represent the map units than others and have been judiciously selected. A statistically valid sampling program was not undertaken due to the cost and time involved relative to the benefits derived.

The biophysical soil landscape legend was developed with first level or broadest stratification based on physiographic regions and associated bedrock geology. The second level of stratification was based on biogeoclimatic zones and subzones. Within this framework the major material types are characterized for texture, drainage, topography, depth, perviousness and soil classification. Table 3 shows the matrix used to construct the biophysical soil landscape legend for the Nechako Plateau and the Bulkley Mountains physiographic regions.

Each biophysical soil landscape unit is a group of soils that have developed on similar parent materials and under broadly similar climatic conditions (in this case, as expressed by biogeoclimatic zones and subzones). These map units provide a sound ecological basis for land use planning and ratings for biological productivity such as ungulate capability ratings (Chapter 7).

A further level of detail can be added to each biophysical soil landscape unit by using map unit modifiers (see the soil map legend, box 4). Map unit modifiers are used to point out areas which differ from the majority of areas, for example areas of wetter soil or areas with steep south aspects.

5.4 BIOPHYSICAL SOIL LANDSCAPES

5.4.1 Introduction

This section summarizes the field observations and describes the character and distribution of the major materials and their associated biophysical soil landscape units (Table 3). Four schematic cross sections show the major characteristics used to differentiate map units (Figures 4 to 7).

Table 3. Biophysical Soil Landscape Units for the Nechako Plateau and Bulkley Range.

Physiographic Region	Nechako Plateau (NE)							Bulkley Range (BR)
Bedrock Grouping	Non-Bedrock Specific			Fine grained Sedimentary and Volcanic			Coarse grained Acid	Fine grained Sedimentary
	SBS(d)	SBS(e)	ESSF	SBS(d)	SBS(e)	ESSF	AT	ESSF
Biogeoclimatic Zone and Sub-Zone*								
Materials Colluvial (deep) (shallow)								
Fluvioglacial		NEG2		NEC2	NEC3 NEC4	NEC5 NEC6	NEC 8	BRC1 BRC1 BRC2 BRC2
Floodplains (gravelly) (sands over gravels) (deep sands)	NEF1	NEF6 NEF3 NEF5						
Fluvial (terraces) (fans)		NEF4 NEF2						
Lacustrine	NEL1							
Morainal (deep) (shallow)		NEM3 NEM4	NEM5 NEM6					
Organic (deep) (shallow)		NEO1 NEO2						
Rock				NER1	NER2	NER3	NER4	BRR1 BRR2

* AT = Alpine Tundra Zone
 ESSF = Engelmann Spruce Subalpine Fir
 SBS(d) = Sub-Boreal Spruce - Spruce Subzone
 SBS(e) = Sub-Boreal Spruce - Subalpine Fir Subzone

5.4.2 Colluvial Terrain and Soils

Shallow colluvial materials are widely distributed within the Lamprey Creek area. Eight biophysical map units were defined and three were sampled. The shallow colluvial areas, map units NEC2, NEC4, and NEC6 are well to rapidly drained. These areas are higher in angular coarse fragments than the deep colluvial areas and are directly derived from the physical weathering of fine grained volcanic and sedimentary bedrock outcrops. The single sample for shallow colluvium had a soil pH of 6.3 for the dark surface organically enriched surface horizon. This lithic south facing site supported a community of kinnikinnick (Arctostaphylos), common juniper (Juniperus communis) and grasses. The organic carbon content of the surface layer was 4.6%.

Map units NEC6, NEC7 and NEC8 were not sampled but are thought to be more deeply weathered and more acidic due to the coarse grained acid intrusive bedrock from which these soils are derived and the higher precipitation in the areas where they occur.

The deeper colluvial materials derived from morainal materials were imperfectly drained (Map Units NEC3 and NEC5). One site had a veneer consisting of 90% coarse fragments over colluviated till. Slope processes are considered active with material moving down the slope, obscuring distinct horizon boundaries. The surface soil pH ranged from 5.4 to 5.9 with relatively high organic carbon content (3 to 4.3 percent). These southerly aspect meadows supported a trembling aspen-saskatoon (Populus tremuloides - Amelanchier alnifolia) vegetation landscape (Chapter 6), explaining the high organic carbon percentages. These sites are considered highly productive and mapped as NEC3 (see Figure 4).

5.4.3 Fluvial Terrain and Soils

Fluvial materials have been deposited by flowing water and materials are often well sorted. The majority of fluvial materials are associated with the major active floodplains in the study area, the Morice, Nanika, Gosnell Rivers and Lamprey Creek. New materials are actively deposited

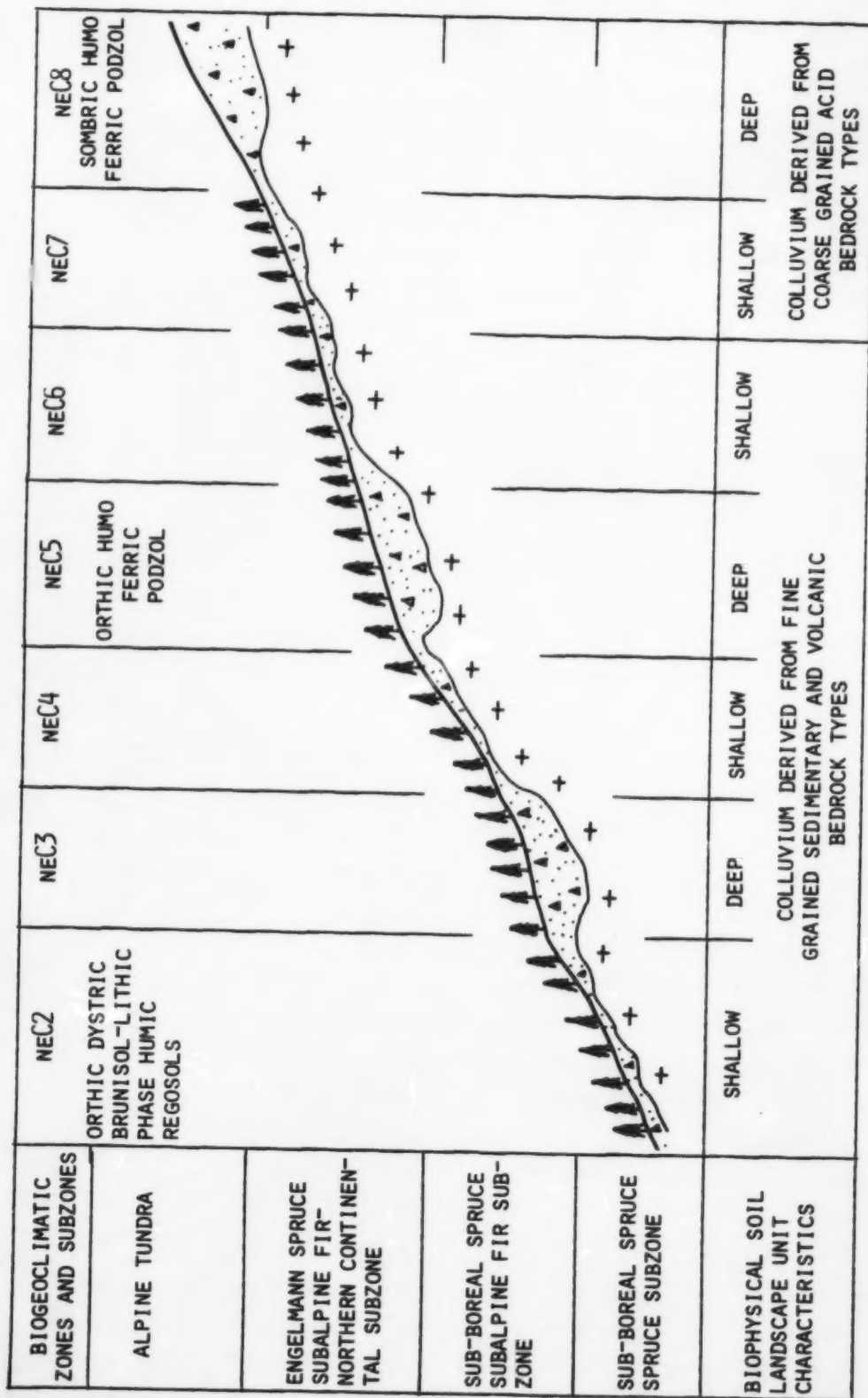


FIGURE 4. SCHEMATIC CROSS-SECTION OF BIOPHYSICAL SOIL LANDSCAPE UNITS FOR COLLUVIAL MATERIALS OF THE NECHAKO PLATEAU.

within these floodplains and channels continue to change. Low terraces not subject to flooding are common along the margins of these floodplains as these watercourses continue to incise into the landscape.

Six biophysical soil map units were defined for fluvial materials within the study area. These fluvial soils can be broadly grouped into 5 types; gravelly textured fluvial terraces (NEF4), sandy floodplains (NEF5), gravelly floodplains (NEF1 and NEF6), gravelly floodplains with a sandy veneer overlying the gravels (NEF3) and gravel fluvial fans (NEF2) (see Figure 5).

Fluvial terraces mapped NEF4 are rapidly drained areas which support xeric plant communities such as the lodgepole pine - lichens vegetation (*Pinus contorta* - *Cladina*) vegetation landscape (Chapter 6). Fluvio-glacial terraces associated with long abandoned meltwater channels have been included with these soils as they have similar soil and plant communities. Coarse fragment contents of these areas are high, (70 to 80%). The three sites used to characterize these terraces had surface soil reaction from pH 4.8 to 5.0, organic carbon of 1.3 percent and extractable iron and aluminum of 1.5 to 2%. Soil matrix textures are coarse and consequently have a low moisture holding capacity.

Deep sandy floodplains NEF5 are associated with lower gradient streams such as the mid section of Lamprey Creek and portions of the Gosnell and Morice river floodplains. The six sites visited ranged from moderately well to imperfectly drained mottled soil indicating a seasonal water table within 0 to 70 cm of the soil surface (NEF5a). Some sandy floodplains were well drained and are mapped as NEF5.

Floodplains characterized by a sandy veneer over gravels are the most extensive (NEF3). The majority of the Morice River floodplain is included within this map unit. Five sites varied in texture from silty loam to loam with soil reaction varying from 3.8 to 6.9 pH.

Frequency of flooding of some portions of the valley floor was not clear and two sites, 82-07624 and 82-07629 had developed Podzolic soil

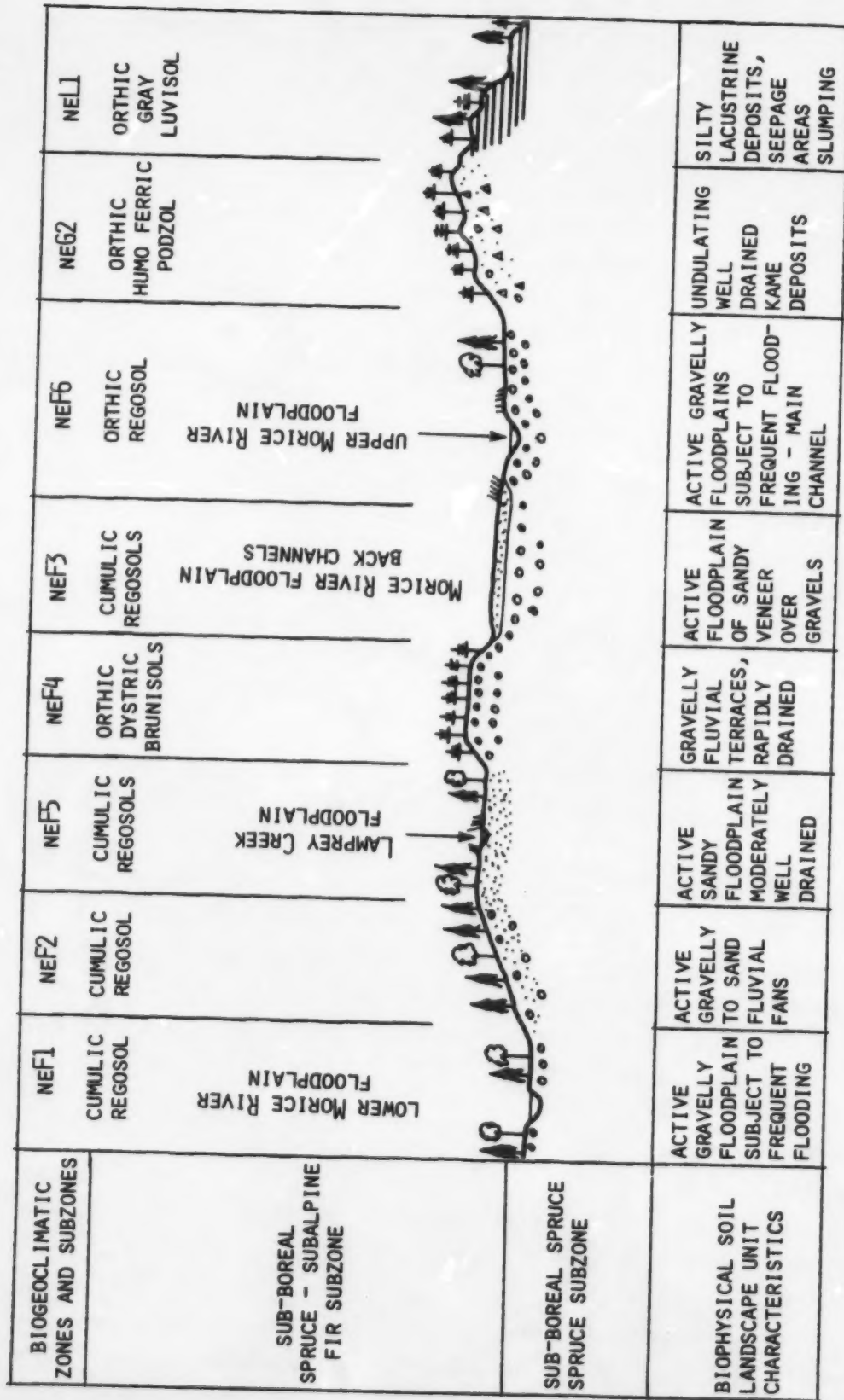


FIGURE 5. SCHEMATIC CROSS-SECTION OF BIOPHYSICAL SOIL LANDSCAPE UNITS FOR FLOODPLAINS, FLUVIAL TERRACES, KAME TERRACES AND LACUSTRINE DEPOSITS OF THE NECKAKO PLATEAU.

profiles with dark brown surface horizons. Podzolic profile development is traditionally believed to take in excess of 100 years and a steady state not reached until 500 years (Chandler, 1942). Only one surface horizon was sampled along the Morice River floodplain in areas which were frequently flooded (pH of 6.9). The two Podzolic profiles had surface pH values of 3.8 and 4.3, indicating a great deal of weathering. Such areas of the valley bottom may not have received deposition of new material for more than 200 years. This does not necessarily indicate that these sites were not seasonally submerged by seepage water, but that such water was not laden with sediments as are the majority of the Cumulic Regosol soils associated with the main channel areas.

A historic look at channel changes has been undertaken by M. Miles (personal communication) through study of 5 sets of aerial photographs available from 1950 to the present. Deep gravelly floodplains (NEF6) are associated with the most active main channels of the Morice, Thautil and Nanika Rivers. These areas are frequently flooded and are sufficiently active that no soil weathering can take place before new material changes the profile.

The lower portions of the Morice River floodplain downstream from Lamprey Creek are climatically warmer, have lower snowfall and are in the Spruce Subzone of the Sub-Boreal Spruce Zone (NEF1).

5.4.4 Fluvioglacial Terrain and Soils

Fluvioglacial deposits are not extensive throughout the Lamprey Creek map area. A major glaciofluvial meltwater channel above the present upper Morice indicates that an abrupt and major change occurred in drainage at one point during deglaciation. This abrupt drainage change left a well established dry river bed some 20 to 30 m above the Gosnell River and along portions of the upper Morice River.

The eastern end of McBride Lake also has evidence of fluvioglacial materials; deep sands, possibly associated with discharge from the Nanika River. The Nanika River channel today dissects fluvioglacial material

along its lower reaches and as it empties into Morice Lake. Kame deposits and discontinuous eskers are also found along Nado Creek and Cedric Creek. Small meltwater channels are scattered within large, deep morainal areas near McBride and Collins lakes.

Soil map unit NEG2 have a wide range of soil matrix textures (<2 mm) ranging from loamy sand, to sand to silt loam and an equally wide range of coarse fragment contents. Most of these soils are well to rapidly drained with undulating to hummocky surface topography.

Two samples with surface pH of 4.2 and 4.7 were taken to characterize map unit NEG2. Both samples had enough extractable iron and aluminum for classification as Podzols.

5.4.5 Glaciolacustrine Terrain and Soils

One small area downstream from the Lamprey Creek and Morice River confluence was identified as glaciolacustrine material (NEL1). A well defined delta, identified with foreset beds exposed in a gravel pit is located in 93L/2 just outside the study area. The glaciolacustrine materials are believed to be associated with this delta and represent the western extent of a much larger lake associated with deglaciation. Ice movement was believed to be from west to east across the map area. Glacial striae on recently exposed bedrock north of Lamprey Lake had an orientation of 90°.

5.4.6 Morainal Terrain and Soils

A majority of the morainal materials are deep and many areas may have a surface layer of less compacted material, believed to be ablation moraine. The matrix textures available for three samples below 2 m in depth were loam and sandy loam. Two other till samples adjacent to the area of Shea Lake and Gosnell Creek also had loam matrix textures.

Soil map units NEM1 and NEM2 occur at the lowest elevations along the Morice River and are associated with the Spruce subzone of the Sub-Boreal

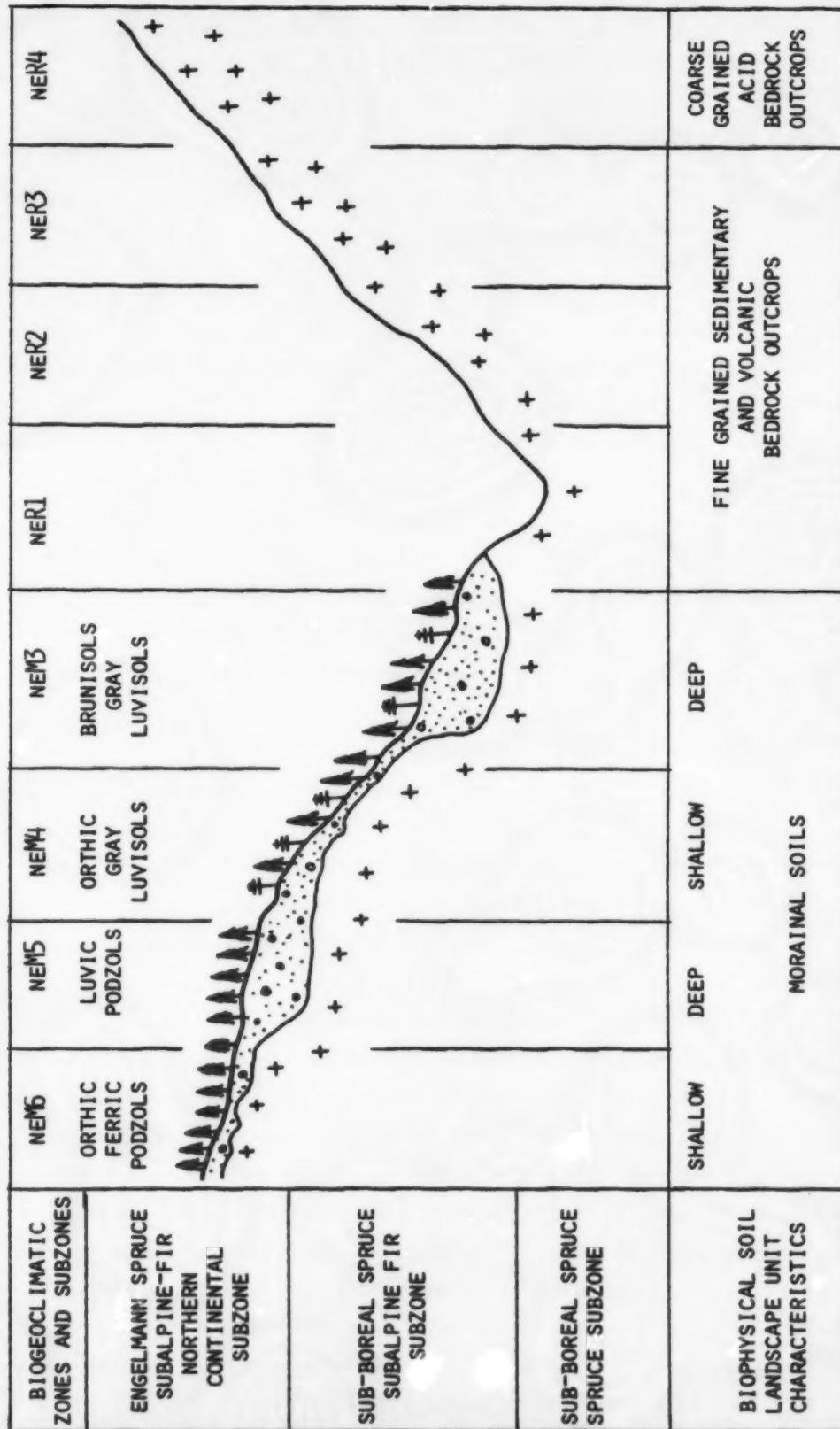


FIGURE 6. SCHEMATIC CROSS-SECTION OF BIOPHYSICAL SOIL LANDSCAPE UNITS FOR MORaine AND ROCK FOR THE NECHAKO PLATEAU.

Spruce biogeoclimatic zone. These map units are an extension of much larger low elevational morainal soils east of the study area and characterized in greater detail by Runka, 1972 (see Deserters Soil Association). Map unit NEM2 is restricted to some south aspects along the lower Morice River valley.

Map unit NEM3 is most wide spread and has six sites available to characterize these areas. Soils mapping has been carried out in the Pimpernel area, mostly detailing drainage characters with NEM3 map units (Lindeburgh and Trowbridge, 1984). NEM3 soils range from 30 to 50% coarse fragments with soil pH from 4.5 to 5.3 for the dark brown podzolic surface horizons. There is evidence, such as clay skins and slightly finer textures, that clay accumulation has taken place below 50 cm in depth. Rooting is above 50 cm with mottles and gleying often associated with the zone of clay accumulation even on moderate slopes. Surface soil textures ranged from loam to clay loam to sandy loam. The medium to coarse soil matrix textures are considered to have been derived from less compact ablation moraine.

Map unit NEM4 is a shallow morainal soil (<1 m). These soils are not as extensive as the deep morainal soils (NEM3) and are commonly mapped as a minor component within a dominately deep morainal area or in association with shallow colluvial materials, (NEC4). The surface pH where sampled was 3.8 with 1.3% carbon and extractable iron and aluminum percentages of 1.6.

Map unit NEM5 is confined to the higher elevations and areas of higher rainfall in the Engelmann Spruce Subalpine Fir Zone. The two sites visited were well and moderately well drained. The upper soil horizons of these sites appeared glacially washed and could have been derived from a layer of ablation moraine overlying more compacted basal moraine. Soil matrix textures (<2 mm particles) were sandy loams, verified by laboratory analysis for both sites. Evidence for clay accumulation was supported by many moderately thick clay skins showing an increase in clay from 9 to 17%. This clay layer is situated from 90 to 120 cm in depth and acts as a layer in which an excess of moisture is present. Common distinct mottles were associated with this layer for the moderately well drained site.

The map unit NEM6 is found in association with deep moraine (NEM5) and shallow colluvial soil (NEC6), usually as a minor component. No sites were visited but soils are considered to have a podzolic development with no significant clay movement in the profile. These soils tend to be well drained on upper slopes, and have a lower moisture holding capacity than the deeper NEM5 soils (see Figure 6).

5.4.7 Organic Terrain and Soils

Organic materials are scattered in distribution and associated with active floodplains and fans. Two organic soil map units (NEO1 and NEO2) have been mapped within the study area. Three organic sites were visited, all were located in poorly drained depressional areas and were greater than 1 metre in depth (NEO1). Map unit NEO2 describes areas of organic deposition of between 40 cm and 1 m in depth (see Figure 7).

5.4.8 Bedrock Terrain and Soils

Two major bedrock groups were recognized: fine grained volcanic/sedimentary bedrock and coarse grained acid bedrock (see Table 2). The most extensive bedrock outcrops are fine grained volcanic and sedimentary. Rock dominated map units often include areas in which <10 cm of material occur over rock as well as bare rock outcrops. The volcanic and sedimentary rocks have weathered and fractured and support xeric vegetation communities (NER1, NER2, NER3).

The acid coarse grained bedrock map units are very limited in distribution within the study area. These areas are also xeric in character with less fracturing of bedrock. Thus, these areas tend to have fewer opportunities for plants to establish (NER4, NER5, and NER6) and tend to be less productive than the fine grained volcanic and sedimentary rock.

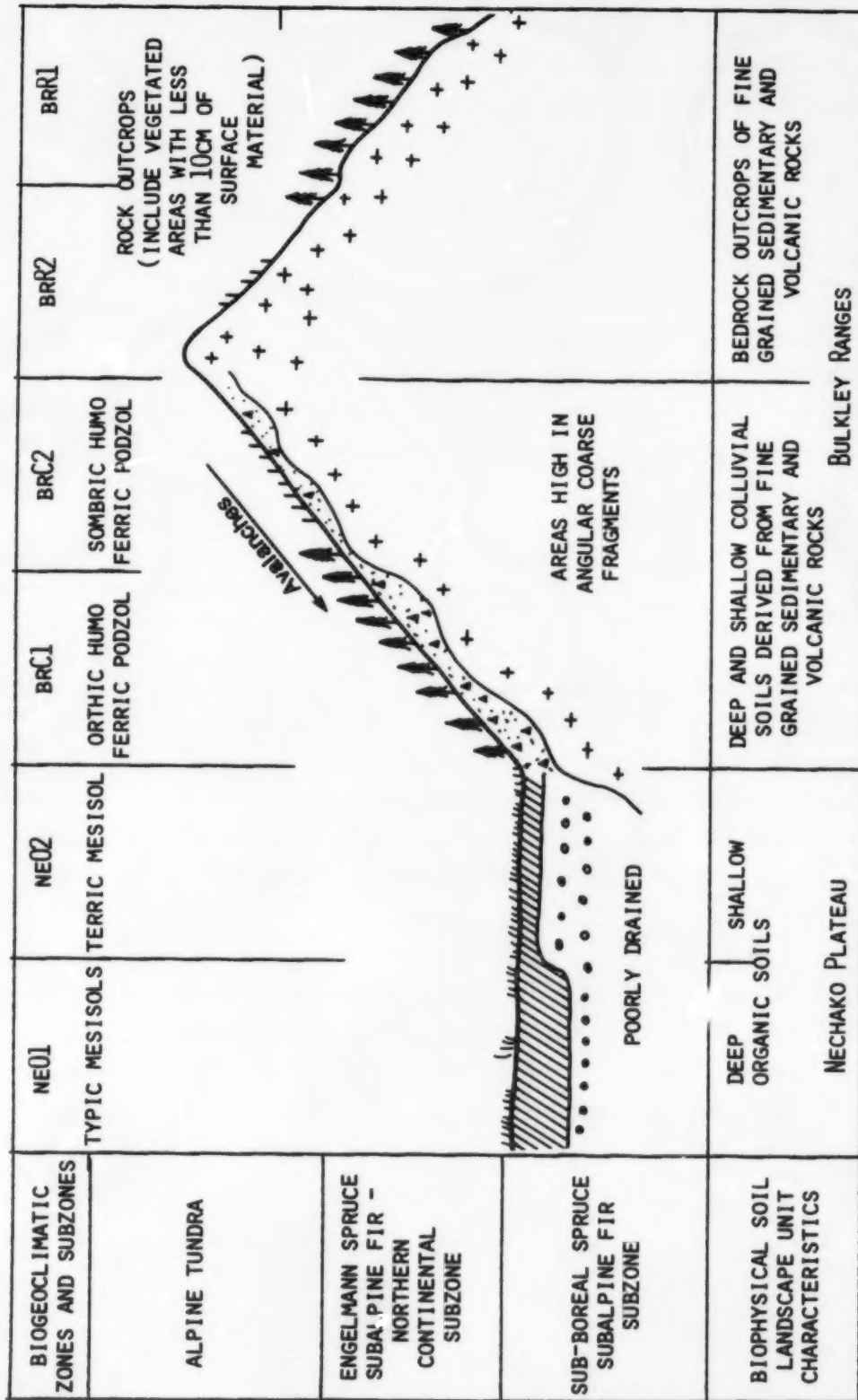


FIGURE 7. SCHEMATIC CROSS-SECTION OF BIOPHYSICAL SOIL LANDSCAPE UNITS FOR ORGANIC AREAS OF THE NECHAKO PLATEAU AND FOR ROCK AND COLLUVIUM IN THE BULKLEY RANGES.

CHAPTER 6. VEGETATION

6.1 VEGETATION ZONATION

The biogeoclimatic zonation of the area is largely described by Pojar, et al. (1984). A draft map at 1:250 000 scale from the Prince Rupert Regional office of Ministry of Forests Research Branch was used as a guide to zonal distribution. The vegetation zonation reflects the physiography and climate of the area. The Coastal Western Hemlock Zone occurs in the western, wetter portion of the study area, while the drier, colder Sub-Boreal Spruce Zone occurs in the rainshadow of the Coast Mountains to the east. Each area has corresponding subalpine and alpine zones also reflecting climatic drying from west to east. This is most notable in the Telkwa Ranges where snow pack is sufficiently reduced by a drier climate and windsweeping to provide good alpine winter range for mountain caribou and mountain goat. The relatively dry and mild climate of the White Spruce Subzone of the Sub-Boreal White Spruce Zone occurs near the lower Morice River. Moose winter range was found to generally be correlated with this subzone, with the exception of strongly edaphic areas.

6.2 VEGETATION LANDSCAPES OF THE MORICE RIVER FLOODPLAIN

The Morice River floodplain study area includes the area downstream of Gosnell Creek on map area 93L/3. It encompasses the transition of the milder White Spruce Subzone (SBSd) to the colder Subalpine Fir Subzone (SBSe) of the Sub-Boreal Spruce Zone (Pojar et al., 1984). This reference contains a detailed description of biogeoclimatic classification in the area.

Vegetation landscapes are areas of similar successional pattern and climax vegetation. The main factors creating these landscapes in the study area are climate, surficial materials, available soil moisture and frequency of flooding. The following is a description of how these factors influence vegetation landscapes in the study area.

- a. Climate. The regional climate of the area (macroclimate) is reflected by limited moisture availability and a shortened growing

season, influencing species diversity and productivity. This regional climate is locally modified by slope, elevation, aspect and cold air drainage. Thus vegetation influenced by a somewhat colder climate occurs along the valley bottom and northerly aspects while a warmer climate prevails on southerly aspects. For example, the trembling aspen - saskatoon vegetation landscape¹ has high browse productivity and low snowdepth, largely attributable to its steeply sloping southern exposure.

- b. Surficial Materials. Surficial materials influence moisture and nutrient availability by their texture and chemical characteristics. For example, relatively dry, nutrient poor ecosystems such as the lodgepole pine - lichens dry forest may develop on coarse textured surficial materials.
- c. Available Soil Moisture. The availability of water to plants can be influenced by soil texture, drainage, soil chemistry and height of water table. Water table height is important in determining the abundance of willows and red-osier dogwood along the Morice River floodplain.
- d. Flooding Frequency. The active floodplain area has had frequent disturbance resulting in a diverse vegetation pattern. This disturbance in addition to water table movement largely contribute to the vegetation landscapes present. These processes along with other biophysical parameters have provided a rich blend of the vital needs of wildlife, i.e. food, shelter and water. This makes the floodplain one of the important wildlife habitats in the area.

The vegetation landscapes of the study area can be grouped into the following landscape systems. They are summarized in Table 4.

- a. Wetland Landscapes. Most wetlands have developed either in depressional upland areas or from a high water table. The former

¹ See vegetation map legend for a further description of these landscapes.

type tend to be organic bogs and fens with lower browse production for moose. The latter type commonly occur at the toe of slopes and near low gradient floodplains. Groundwater movement in these areas usually is a major contributor to their high browse productivity.

- b. Riparian Landscapes. These landscapes are dominantly influenced by the hydrology and erosional cycles of the river. Evidence of flooding in soil profiles was found on most areas of the valley bottom. Primary succession (pioneer) is the dominant successional process.
- c. Fluvial Terrace Landscapes. These areas are no longer influenced by flooding. Parent materials have a strong influence on moisture holding capacity and nutrient availability. Some sandy gravelly terraces can be very dry and nutrient poor, as in the lodgepole pine - lichens vegetation landscape.
- d. Upland Landscapes. These areas may develop climatic climax forests through a series of seral stages (secondary succession). These stages may include either lodgepole pine or trembling aspen forests. Browse productivity is usually higher on the latter type, being more common on steeper south aspects.
- e. Rock Outcrop Landscapes. These areas have low vegetative productivity because of their dryness and thin soil. Steep south aspects usually provide some mule deer range.

Table 4. A Summary of the Vegetation Landscapes of the Morice Floodplain Study Area*.

Landscape System	Map Symbol	Vegetation Landscape Description
Wetland Landscapes	SB	Sedge - bog birch - sphagnum wetlands on organic soils.
	WS	Wetland willow - spiraea shrub on gleysols.
Riparian Landscapes	AW	Fluvial alder - willow shrub.
	CW	Black cottonwood - willow deciduous forest.
	SS1	Black cottonwood - red-osier dogwood deciduous forest.
	SS2	Black cottonwood - white spruce mixed forest.
	SS	White spruce coniferous forest.
Fluvial Terrace Landscapes	LL	Lodgepole pine - lichens dry forest.
Upland Landscapes	AS	Trembling aspen - saskatoon.
	BS1	Trembling aspen - white spruce seral forest.
	BS2	Lodgepole pine - white spruce seral forest.
	BS	White spruce - alpine fir coniferous forest.
Rock Outcrop Landscapes	R	Rock Outcrops.

*See the vegetation landscape map for a full description of these types.

CHAPTER 7. WILDLIFE CAPABILITY

7.1 INTRODUCTION

The four wildlife maps depict capability for moose, mountain goat, woodland caribou and mule deer. Capability is largely reflected by present use, mainly because access to much of the area is limited. Exceptions to this include the caribou of the Telkwa Range, which have been recovering from overhunting of the 1940's and 1960's (Hodson, 1977); and moose in the Lamprey Creek map area which appear to have not yet responded to the increase in forage created by logging. Despite these exceptions the maps should provide a good basis for survey stratification, planning and habitat protection.

Greater survey effort was directed towards areas of higher ungulate capability. These include portions of the Telkwa Ranges with high capability for goat and caribou, and the lower Morice area with high moose capability. The possible detrimental effects of the proposed Kemano Completion hydroelectric development on moose habitat along the Morice River increased our survey effort in this area. The mountainous western half of the study area was surveyed with the least intensity because of its remoteness and frequent poor weather. Survey reliability is expected to vary accordingly.

7.2 BACKGROUND WILDLIFE SURVEYS

The limited background information pertaining to wildlife in the study area is listed here for reference:

1. "A fisheries and wildlife survey of the Burnie Lakes park proposal", an unpublished Parks Branch document, dated 1975.
2. "Snowmobile conflict with the Telkwa caribou herd," an unpublished B.C. Fish and Wildlife Branch document, dated 1980.

3. "A winter population and habitat survey in the Telkwa Mountains of mountain caribou, January - April/77," an unpublished B.C. Fish and Wildlife document.
4. "Gosnell resource plan," an unpublished document by Northwood Pulp and Timber Limited, undated.
5. B.C. Fish and Wildlife Branch regional files, Smithers.
6. Canada Land Inventory winter flights of 1968 and 1974, and 1:250,000 scale capability map.

7.3 METHODS

Background data pertaining to wildlife and biophysical themes were used in combination with aerial photographs to prestratify the sampling prior to field work. Summer field work (1983) was done concurrently with the soils, terrain and vegetation sampling. Fifteen days of summer field work and 16 hours of helicopter time were expended, mainly to gain ground access to selected sites, although one flight focused on mountain goat habitat use in the Telkwa Ranges (September 19, 1983). Data from ground stops was recorded using the procedures described in Walmsley et al., 1980. The data from all reconnaissance plots (site, soil, vegetation and wildlife forms) has been entered into the B.C. Soil Information System. Ministry of Forests Research Branch personnel provided valuable field and office assistance.

Winter field work was mainly by helicopter to confirm ungulate use and habitat conditions of suspected winter ranges. Four days and twelve hours of helicopter time were expended between January 31 and February 6, 1984. The flight lines and observations were recorded using the methods described in Demarchi et al., 1983. These and the flight records of the summer field work are available from the Habitat Management and Inventory Section of the Wildlife Branch, Ministry of Environment, Victoria, B.C. Wildlife Branch personnel from Smithers assisted with the winter flights.

The mapping procedure follows the methods outlined in "Wildlife Capability Classification for British Columbia: An Ecological (Biophysical) Approach for Ungulates." Briefly, the biophysical data or maps of terrain, soils, climate and vegetation are grouped or subdivided into units of similar ecological significance to ungulates. These units are separated into either summer or winter range and ranked according to their perceived importance. This ranking is then placed in the context of provincially based carrying capacities described by Demarchi et al., 1983. Census information is used to support these carrying capacity estimates where available.

7.4 CAPABILITY ASSESSMENT

7.4.1 Moose

Moose are the most abundant and widespread of the ungulates present in the area during the summer. Important summer areas identified include the floodplains and associated wetlands of the major watercourses in the area, notably the Morice, Thautil, Burnie, upper Clore and Atna Rivers, as well as Gosnell and Lamprey Creeks. These areas provide calving and rutting areas in addition to acting as corridors for movement. The lower Clore River appeared to be geographically isolated from moose populations east of the Coast Mountains.

Winter habitat use by moose is low in most of the study area because of deep snow. High winter use occurs along the Morice River floodplain from the area of the Lamprey Creek confluence and downstream. The most heavily used habitats appear to be the alder-willow river bars and adjacent forest. These habitats are maintained by flooding and provide abundant forage combined with some snow interception. The forage productivity and lower snow depth of this area largely corresponds with the limit of the SBS(d) subzone (Pojar et al., 1984). Moose habitat along the Morice floodplain is discussed in greater detail in Section 7.5. The Lamprey Creek floodplain also provides high capability winter range downstream of Pimpernel Creek. The extensive logging in the Lamprey Creek map area has resulted in abundant willow regeneration on upland areas. Willow appears to be more abundant at

lower elevations (below 1,000 m elevation or slightly higher on south aspects). Such areas may provide good early winter range, but are considered to have excessive snow depths on average winters. An area of moderate capability winter range was identified along the south aspects of Tagetochlain Lake. This area has high present forage production but relatively deep snow, although snow depth is somewhat reduced on southern exposures.

7.4.2 Mountain Goat

Mountain goat is the most abundant ungulate species in the study area during the winter months. Their winter range is limited to areas where wind-sweeping and insolation expose forage and alter snow condition to allow movement. Such areas are usually convex in shape to be windswept and reduce avalanche hazard. Areas which were exposed to avalanche hazard from above but which otherwise appeared suitable were seldom observed to be used. With the exception of the Telkwa Ranges, all winter range was mapped at or below krummholz elevation, mainly because of the lack of vegetation at higher elevations as a result of extreme climate and persistent snow cover. The occurrence and extent of good winter range is therefore mainly the fortuitous occurrence of favourable landform at the appropriate aspect and elevation in combination with climate. Such areas are more restricted in area and less common in the western half of the study area because of deeper and more persistent snow cover. Goat winter range was mapped at lower elevations here (below 1525 m) because the upper elevations of krummholz are lower, and subalpine fir (suspected to be a major winter forage species) is more common in the subalpine forest than in the mountain hemlock dominated krummholz. Goats were observed in steep rock-bluff subalpine forest in several locations especially on the Howson Range and west of the Clore and Burnie Rivers. Areas of high capability were identified here, but they tended to be smaller than areas east of the Bulkley Ranges. In addition, the alpine zone tends to have more vegetation than areas to the west. Goats were commonly seen in the alpine between 1800 and 2000 meters.

Goat ranges which appeared transitional between the ecotype found in the Telkwa Range and that of the high snowfall areas, occur in the area north-

east of Mt. Loring and on Herd Dome. There appears to be some rainshadow effect here, although not as great as in the Telkwa Range.

The spectacular canyon of the Clore River immediately downstream of its confluence with the Burnie River appeared to provide good year-round goat habitat. Difficulty in spotting goats because of forested cover, no helicopter landing sites and probable movement between there and the adjacent Hope and Pillar Peaks all combined to make a capability assessment of the canyon difficult.

During the summer months goats range widely throughout the mountainous portions of the study area, often to higher elevations and on aspects not used during the winter. Little specific is known about important summer habitat in the Bulkley Ranges.

7.4.3 Woodland Caribou

Caribou appear widespread during the summer months, with records from the Burnie Lakes area, Herd Dome and the Mt. Loring area. Both valley bottom and rolling alpine habitat were identified as summer range in this area, although this was not confirmed by observation of present use. Old antlers and historical records occur throughout these areas. Caribou capability was not mapped west of the Burnie/Clore confluence, as caribou are seldom reported from coastal forest, and steep slopes often form a barrier to movement. The rolling kummholz and subalpine areas to the east of Burnie Lakes appears to be more frequently used summer range, probably because of its landform and proximity to the Telkwa Range wintering area.

The winter range distribution of caribou appear to focus on the rolling alpine areas of the Telkwa Range. These areas have the convex landform, lower snow depth, windsweeping and available vegetation necessary to support caribou during the winter. Several plots recorded an abundance of terrestrial lichens, grasses and sedges during both summer and winter. Such alpine habitat surround the rugged portion of the Telkwa Range, with some areas occurring to the north of the study area. Large areas occur in the upper Houston Tommy and Emerson Creek drainages. These more extensive areas are suspected to be important in providing a variety of areas with available

forage, depending on snow condition and windsweeping. Disturbance of caribou on these areas by snowmobiles or other activity (snowmobile tracks were observed throughout the Emerson/Houston Tommy areas on February 4, 1984) could shift use to less suitable areas or cause excessive use of less disturbed areas. The detrimental effects of such activity would be difficult to measure, other than by a long-term decline in herd numbers.

Use of winter habitat other than alpine probably occurs during early winter or during periods when alpine forage is unavailable. These areas are not well known and have largely not been identified on the capability maps. Mature subalpine forest with a high arboreal lichen density may provide winter range at some times. Such areas typically have a southerly aspect and are seldom steeply sloping. Forested areas with high terrestrial lichen ground cover are also commonly used by caribou during the winter, although this was unconfirmed in the study area. This would likely be an early winter activity as snow depths in most of the area are excessive for such foraging by mid-winter. Some of these areas were identified and include fluvioglacial deposits near Mooseskin Johnny Lake, along the Thautil and upper Morice Rivers, and along lower Houston Tommy Creek.

7.4.4 Mule deer

Mule deer range widely throughout the lower elevations of the eastern half of the study area during the summer months. They appear to occur infrequently in the western half of the study area, presumably because of its distance from winter range. Steep south aspects and clearcuts appeared to be most frequently used.

Winter range is very restricted in the study area because of deep snow. Deer were observed on February 6, 1984 on the south aspects of the Morice River near Lamprey Creek, although snow depths were below normal at low elevation. Two other areas of low to moderate capability deer habitat occur on the south aspects of lower Houston Tommy Creek and Tagetochlain Lake, although both areas usually have an excessive snow depth for deer and were not considered to be winter range.

7.4.5 Other species

While habitat for other species are not mapped, some habitat generalizations can be made.

Grizzly bear frequent most of the mountainous portion of the study area. Some of the more important areas of habitat can be identified using the environmental condition parameters on the ungulate maps. Some of these include: rolling alpine or krummholz areas (EaLr or EkLr), avalanche areas (La), subalpine meadows (Em), seepage areas (Eh), organic areas (Eo) and floodplains (Lf). Important spring ranges likely occur on main floodplains and south-facing avalanche tracks.

Black bear use was most commonly noted (July, 1983) in the clearcut logging of the 93L/3 map area and along major floodplains throughout the study area. Numerous black bears appeared to be foraging on crowberry (Empetrum nigrum) at krummholz elevation during the mid-September field survey.

Four trumpeter swans were noted on the Burnie River near the outlet of Burnie Lake on February 4, 1984. The river was unfrozen in this area and may provide a common wintering area.

7.5 IMPLICATIONS OF FLOW REDUCTION ON MORICE RIVER MOOSE HABITAT

The Kemano Completion Project would result in an average flow reduction of 20% at Owen Creek with summer flows reduced 30-40% (Department of Fisheries and Oceans, 1984). The degree to which this affects moose habitat along the Morice River floodplain depends on the amount of flow reduction, the degree of freshet flooding and the extent of habitats dependent on disturbance and/or a high water table. The probable effects are discussed here in the context of the vegetation landscapes described in Chapter 6. This discussion includes only the portion of the Morice River floodplain in the study area and does not include winter range downstream. Larger areas of winter range floodplain occur downstream (Stewart, 1981) and would be similarly affected.

The fluvial alder-willow (AW), black cottonwood-red osier dogwood deciduous forest (SS₁), and black cottonwood-willow deciduous forest (CW) vegetation landscapes have high forage production. They constitute the recently disturbed portions of the floodplain. The black cottonwood - white spruce mixed forest (SS₂) and white spruce coniferous forest (SS) vegetation landscapes are seral to SS₁ and produce progressively less forage with time as the cover of spruce forest increases. These older stands have moderate to high snow interception potential and function mainly as cover. Flooding, resulting in overbank deposition, channel erosion and aggradation, functions to maintain a dynamic balance of the forage and cover producing landscapes. Elevation and movement of subsurface water in these stands contributes to the growth of red osier dogwood and willow, probably the most important browse species on the floodplain.

A reduction of flow would likely have several affects on succession. The AW vegetation landscape would succeed to CW, CW would succeed to SS₁ and the SS₁ or SS₂ landscapes would succeed to their SS spruce climax more rapidly. The areas of habitat affected were not calculated because the mapping scale (1:50 000) does not show sufficient detail and most winter range lies downstream. However, considerable long term (50 years plus) loss of forage production and associated carrying capacity would likely occur. Coniferous landscapes (SS) with lower forage production would become more abundant. There would be initial increase or maintenance of the present area of winter range while pioneer vegetation establishes on exposed river bars. However, it is likely that the quality of much of this range will decline even in the short term because of the reduced elevation and movement of subsurface water.

Summer use could also be affected. Moose commonly use islands and floodplain areas for calving. A reduction in island security and dense back channel vegetation would reduce the suitability of the area for calving and summer use.

CHAPTER 8. SUMMARY

The Morice Biophysical Study provides reconnaissance level terrain and ungulate capability maps for four 1:50 000 scale map areas (93L/3-6). In addition, soil landscapes are provided for the Lamprey Creek area (93L/3) and vegetation landscapes for the Morice River floodplain (portion of 93L/3). The maps and report provide a stratification of these resources and themes for resource planning and management.

The physiography of the area includes the rugged Coast Mountains along the western margin of the study area, the central Hazelton Mountains and the subdued Nechako Plateau to the southeast. The climate of the area is strongly affected by this physiography. The westerly prevailing winds lose much of their moisture in the Coast Range. Areas to the east, notably the Telkwa Range and lower Morice River valley have high capability wildlife habitats largely attributable to this combination of physiography and lower snow depth.

The terrain of the area largely reflects its glacial history and is one of the dominant features affecting ungulate habitat. Glacial till is common on lower valley slopes and on the Nechako Plateau. It provides some of the better upland moose habitat because of its higher moisture and nutrient status, such as in the Lamprey Creek map area. Colluvial materials dominate the slopes of mountainous areas. Their higher elevation usually results in deep winter snow cover, except where windsweeping occurs. Glaciofluvial materials are common in major valley bottoms. Their very coarse texture, dryness and low nutrient status all contribute to low vegetative productivity. However, such sites may have abundant terrestrial lichen cover suitable for woodland caribou forage such as along Houston Tommy Creek and the Thautil River. Active fluvial materials are most common along larger, lower gradient watercourses such as the Morice and Clore Rivers. The seasonally high soil moisture content and frequent flooding of these areas contributes the high forage production of to this habitat. Organic materials are common in the depressions of the hummocky mantle of till on the Nechako Plateau. Many of these areas provide important summer range for moose.

The biophysical soil landscapes for the Lamprey Creek map area are based on a stratification of physiographic regions, bedrock geology, biogeoclimatic zonation, topography, terrain and soil properties. Each soil landscape type so derived has similar ecological properties and provides much of the basis for the ungulate capability assessment of this area. In all 40 soil landscapes were described in two physiographic regions and four biogeoclimatic subzones.

The Morice River floodplain was studied in greater detail to investigate the potential impact of flow reduction on moose habitat by the Kemano Completion Project. Thirteen vegetation landscapes describe areas of similar vegetation and succession in this area. A reduction in flow would likely reduce the creation of riparian habitat on floodplain, and reduce the productivity of existing habitat over the long term.

Ungulate capability is strongly influenced by the physiography and climate of the area. Deep snow results in most of the area being limited to summer range use however, very high capability for moose, woodland caribou and mountain goat exists where this deep snow is reduced by local factors. The restricted extent of such areas emphasizes the importance of this habitat. In the western portion of the area, only mountain goats have adequate habitat to remain in the area through the winter. Moose winter range occurs along the Morice River floodplain and other nearby areas of combined reduced snow depth and high forage abundance. The Telkwa Range provides areas of very high capability mountain goat and woodland caribou winter habitat as a result of its reduced snowfall and favourable topography. This area of spectacular rolling alpine in combination with rugged cliffs makes it an area of provincial significance for wildlife habitat.

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APPENDIX A

Distribution of Plots in the Biophysical Soil Landscapes
of the Lamprey Creek Map Area, 93/L 3

Soil Landscape Map Unit	93/L 3 Data Form Numbers	Adjacent Map Areas 93/L 4 & 6
NEC2	82-07620, 82-07608, 78-1628 82-07607 82-07615	
NEC3		
NEC4		
NEC5		
NEC6		
NEC7		
NEC8		
NEF1		
NEF2		
NEF3	82-07622, 82-07624, 82-07625, 82-07626 82-07629	
NEF4	82-07610, 78-1636, 78-1626, 78-1618	
NEF5	82-07601, 82-07612, 82-07613, 82-07621 82-07623, 82-07630, 82-07631, 82-07627	
NEF6	82-07628	
NEG2	82-07619, 8207632	
NEL1		
NEM3	82-07603, 82-07605, 82-7614, 82-07616 82-07616, 82-07618, 78-1629	82-17752, 82-17754 82-17783
NEM4	82-07606	82-17756
NEM5	82-07617, 82-07611, 82-07604	
NEM6		
NEO1	82-07617, 82-07611, 82-07604	82-17751
NEO2		
NER1		
NER2		
NER3		
NER4		
BRC1		
BRC2		
BRR1		
BRR2		

APPENDIX B

Sources of further information

1. Data collected during field studies is available by contacting the authors. This includes:
 - plot data including site, soil, terrain, vegetation and wildlife forms (Walmsley et al., 1980).
 - laboratory analyses of soil samples.
 - ungulate habitat aerial survey flight lines and transcripts.
 - background biophysical information.

2. Ungulate capability maps (1:50 000 scale) bordering this study area on 93L only, their accompanying report: A Moose Habitat Assessment of the Bulkley - Endako Area (Fuhr and Pendergast, 1983), and copies of this report are available from:

Maps B.C.
Surveys and Resource Mapping Branch
553 Superior Street
Victoria, B.C.
V8V 1X5

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